

# 1 Stand Age and Climate Change Effects on Carbon 2 Increments and Stock Dynamics

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14 **Abstract:** Carbon assimilation and wood production are influenced by environmental  
15 conditions and endogenous factors, such as species auto-ecology, age, and hierarchical  
16 position within the forest structure. Disentangling the intricate relationships between those  
17 factors is more pressing than ever before due to the pressure of climate change. Yet, our  
18 understanding of how future climate will interact with forests of different ages is  
19 particularly limited, and only a few studies have explored this relationship under changing  
20 climate conditions.

22 We employed a validated process-based forest model for simulating undisturbed forests of  
23 different ages under four climate change scenarios (plus one no climate change) coming  
24 from five Earth System Models. In this context, carbon stocks and increment were  
25 simulated via total carbon woody stocks ( $MgC\ ha^{-1}$ ) and the mean annual increment ( $m^3\ ha^{-1}year^{-1}$ ), which depend mainly on age and long-term processes, such as climate trends.  
26 We find greater differences among different age cohorts under the same scenario than in  
27 different climate scenarios under the same age class. We found different C-accumulation  
28 patterns under climate change between coniferous stands and broadleaves. Increasing  
29 temperature and changes in precipitation patterns led to a decline in above-ground biomass  
30 in spruce stands, especially in the older age classes. On the contrary, the results show that  
31 beech forests at DK-Sor will maintain and even increase C-storage rates under most RCP  
32 scenarios. Scots pine forests show an intermediate behavior with a stable stock capacity

34 over time and in different scenarios but with decreasing mean volume annual increment.  
35 These results confirm current observations worldwide that indicate a stronger climate-  
36 related decline in conifers forests than in broadleaves. We, therefore, advocate for a better  
37 understanding of the interaction between forests and climate to better inform forest  
38 management strategies, ultimately dampening the impacts of climate change on forest  
39 ecosystems.

40

41 **Keywords:** carbon cycle, climate change, forest age, forest management, carbon stocks

42

### 43 **Introduction**

44 Assessing the quantity of CO<sub>2</sub> equivalent stored in forest ecosystems is one of the main  
45 goals for implementing the new European Forest Strategy for 2030, a key component of the  
46 European Green Deal, to achieve greenhouse gas emission neutrality by 2050. Within this  
47 framework, European forest strategies have been geared towards forest-based mitigation  
48 plans [1, 2, 3], which makes it essential to estimate the carbon sequestration capacity and  
49 potential under future climate conditions.

50 In the near future, Europe and Mediterranean areas will emerge as focal points ('hot spots')  
51 of climate change, characterized by heightened temperatures and environmental impacts [4,  
52 5]. Carbon assimilation and wood production are influenced by environmental conditions  
53 (e.g., precipitation, temperature, atmospheric CO<sub>2</sub>, etc.) and endogenous factors, such as  
54 species auto-ecology, age, and hierarchical position within the forest structure. In the past  
55 decades, forest ecosystems proved to be crucial net carbon sinks [6, 7], likely due to the  
56 positive fertilization effects of rising atmospheric CO<sub>2</sub> and temperature [8]. However,  
57 whether this effect will remain positive or be compensated by other limiting factors is still  
58 a matter of debate [9, 10, 11]. Some studies suggest that the fertilization effect on carbon  
59 storage and biomass production fades with forest aging in temperate forests [12, 13] since  
60 these positive effects cannot continue indefinitely, complicating the picture of the forest  
61 response to climate changes even further. This is already the case in Europe, where forest  
62 aging and increased disturbances are causing the saturation and decline of the forest carbon  
63 sink [10]. Unfortunately, there is not yet a clear strategy to increase the mitigation  
64 potentials of forests, and the factors involved are manifold and entangled together [12, 14,  
65 15].

66 The need to disentangle the intricate relationships between those factors is even more  
67 pressing under climate change. Our current understanding of how future climate will

68 interact with forests of different age classes is particularly limited, especially since only a  
69 few studies have explored the relationship between age and the ecosystem's carbon balance  
70 under changing climate conditions [16].

71 The climate sensitivity of age cohorts is driven, among all, by different access to  
72 environmental resources, such as root depth and, therefore, access to water, as well as  
73 height, which affects leaf-level water potential and, thus, stomatal conductance [17].  
74 Rooting depth and height jointly affect the tree's sensitivity to water scarcity, a key  
75 environmental driver of change. Future changes in environmental conditions are expected  
76 to impact the age spectrum differently [18, 19, 20].

77 Since forest age is determined by management practices and 75% of European forests are  
78 even-aged [21, 22], it is crucial to grasp and pin down the role of age in the sensitivity of  
79 forest carbon stocks to climate change to guide and inform adaptative forest management.  
80 Process-based forest models enable the exploration of climate change impacts on various  
81 age cohorts within the same area, a task difficult to achieve through direct field  
82 measurements, which would require decades or more. In this regard, this study examines  
83 the ability of different forest age classes under the same future climate conditions to sustain  
84 high productivity and carbon stock capacity. To achieve this goal, we employed the 'Three  
85 Dimensional - Coupled Model Carbon Cycle - Forest Ecosystem Module' (3D-CMCC-  
86 FEM) [23, 24], simulating undisturbed forests of different cohorts under four climate  
87 change scenarios (and including one 'no climate change' scenario), from the moderate one  
88 (RCP 2.6) up to the most severe one (RCP 8.5) coming from five Earth System Models. In  
89 this context, carbon stocks and increment were simulated via total carbon woody stocks  
90 (TCWS, i.e., the standing woody biomass in  $MgC\ ha^{-1}$ ) and the mean annual increment  
91 (MAI, in  $m^3\ ha^{-1}\ year^{-1}$ ), which depend mainly on age and long-term processes, such as  
92 climate trends.

93 The primary aim of this research is to explore (i) the direct effects of climate change on the  
94 overall carbon storage capacity across various stands, species, and age classes situated in  
95 diverse regions of Europe; (ii) elucidate the potential influence of forest age on stand  
96 dynamics in adapting to forthcoming climate shifts.

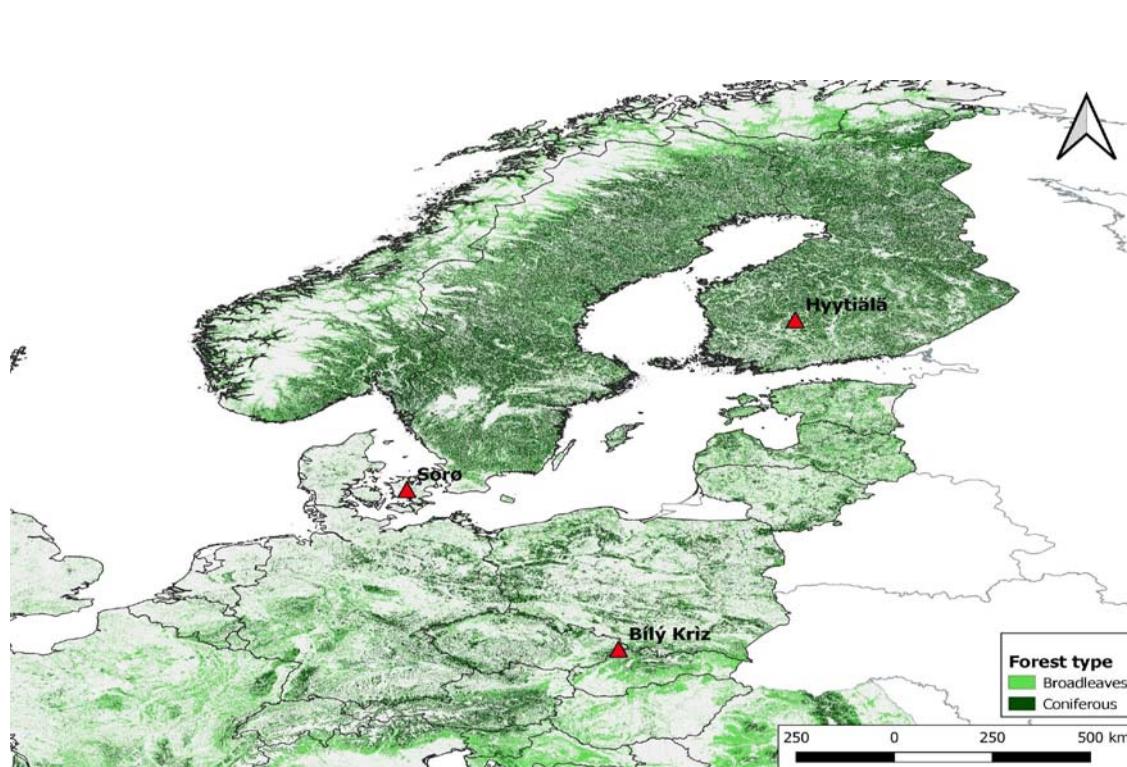
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## 98 2. Materials and Methods

### 99 2.1. Study sites and virtual stands

100 The study was conducted in three even-aged, previously managed European forest stands i)  
101 the Boreal Scots pine (*Pinus sylvestris* L.) forest of Hyytiälä, Finland (FI-Hyy); ii) the wet

102 temperate continental Norway spruce (*Picea abies* (L.) H. Karst) forest of Bílý Kríz in the  
103 Czech Republic (CZ-BK1); and iii) the temperate oceanic European beech (*Fagus sylvatica*  
104 L.) forest of Sorø, Denmark (DK-Sor) where the 3D-CMCC-FEM (in different versions)  
105 has been already validated in the past [15, 25, 26]. For each site, daily bias-adjusted  
106 downscaled climate data from five Earth System Models (i.e., HadGEM2-ES, IPSL-  
107 CM5A-LR, MIROC-ESM-CHEM, GFDL-ESM2M, and NorESM1-M) driven by four  
108 Representative Concentration Pathways, namely RCP 2.6, 4.5, 6.0, and 8.5 were available  
109 [27, 28] (Fig. S1). For more detailed information on the study site characteristics and  
110 climate data, see [15, 25, 26, 29]. The chosen sites have been selected due to their long  
111 monitoring history and the availability of a wide range of data sources for both carbon  
112 fluxes and biometric data for model evaluation, as well as bias-corrected climate scenarios  
113 for simulations under climate change scenarios from the ISIMIP-PROFOUND initiatives  
114 (<https://www.isimip.org/>)[26, 29]. In addition, these stands: i) represent the most common  
115 European tree species; ii) their current state is the result of the legacy of past forest  
116 management; iii) they are mainly mono-specific and therefore represent interesting «living  
117 labs» to study the effects of climate change on single-species and their productivity,  
118 reducing confounding effects which otherwise make models struggle to predict forest  
119 growth and carbon dynamics (e.g., [30, 31]), iv), and they have already been investigated  
120 in the context of climate-smart-forestry silvicultural scenarios [15].



123 **Figure 1.** Test site locations in Europe.

124

125 **2.2. The model**

126 The ‘*Three Dimensional - Coupled Model Carbon Cycle - Forest Ecosystem Module*’ (3D-  
127 CMCC-FEM v 5.6 [13, 15, 23, 24, 25, 32, 33] is a biogeochemical, biophysical, process-  
128 based, stand-level forest model. The model is built to simulate carbon, nitrogen, and water  
129 cycles in forest ecosystems, even including forest dynamics, under scenarios of climate  
130 change and disturbances (e.g., forest management) and parameterized at the species level.  
131 Photosynthesis is modeled through the biogeochemical model of Farquhar von Caemmerer  
132 and Berry [34] implemented for sun and shaded leaves [35] (de Pury and Farquhar, 1997)  
133 and parameterized as in Bernacchi et al. [36, 37]. Temperature acclimation of leaf  
134 photosynthesis to increasing temperature is accounted for following Kattge and Knorr [38].  
135 Autotrophic respiration ( $R_A$ ) is modeled mechanistically by distinguishing the costs of  
136 maintaining already existing tissues ( $R_M$ ) and the cost of synthesizing new ones ( $R_G$ ).  
137 Maintenance respiration is controlled by the amount of nitrogen (stoichiometrically fixed  
138 fraction of live tissues) and temperature. Temperature effects on enzyme kinetics are  
139 modeled through a standard Arrhenius relationship but acclimated for temperature as  
140 described in Collalti et al. [25]. The net primary productivity (NPP) is the gross primary  
141 productivity (GPP) less  $R_A$ . Not all the annual NPP goes for biomass production since the  
142 model considers the Non-structural carbon (NSC) pool, an additional seventh C-pool which  
143 includes starch and sugars (undistinguished) used to buffer periods of negative carbon  
144 balance (when respiration exceeds assimilation; i.e.,  $R_A > GPP$ ). Ultimately, the more trees  
145 respire, the more NSC is used to sustain metabolism and NSC pool replenishment, and the  
146 less NPP and BP there are (and less carbon is stocked). In the extreme case, when and if all  
147 NSCs are depleted because of metabolism without being replenished through current  
148 photosynthates, the model predicts stand mortality based on the carbon starvation  
149 hypothesis [39, 40].

150 The phenological and allocation schemes are all described extensively in Collalti et al. [23,  
151 24, 41] and Merganičová et al. [42]. The 3D-CMCC-FEM accounts for the ‘age-effect’ in  
152 several ways. ’60s ecological theories describe [43, 44], and past and growing pieces of  
153 evidence suggest that stabilization and a further slight decline follow an initial step-wise  
154 increase in forest productivity. The causes of such a decline are debated and include a  
155 decline in the GPP because of hydraulic limitation [17, 45] or an increase in  $R_A$  because of  
156 increased respiring biomass [19, 20, 46]. The 3D-CMCC-FEM accounts for both by  
157 including an age modifier [47], which reduces maximum stomatal conductance (and then

158 also GPP) in the Jarvis model and increases  $R_A$  because of biomass accumulation during  
159 forest development.

160

161 *2.3. The model runs and results evaluation.*

162 The 3D-CMCC-FEM was first evaluated under observed climate and field data for GPP  
163 and  $NPP_{woody}$  (i.e., the NPP for woody compound;  $gC\ m^{-2}\ year^{-1}$ ) and the diameter at breast  
164 height (DBH)(see ‘Model validation’ paragraph in Supplementary Material; [13, 15]). The  
165 model was forced with the modeled climate under different emission scenarios,  
166 corresponding to the RCP atmospheric  $CO_2$  concentration values for the period 1997 to  
167 2100, ranging from  $421.4\ \mu mol\ mol^{-1}$  in the ‘best-case scenario’ (RCP2.6) to  $926.6\ \mu mol\ mol^{-1}$   
168 of the ‘worst-case scenario’ (RCP 8.5) coming from the ISIMIP-PROFOUND  
169 initiative. For comparison purposes, we forced the forest model with a detrended and  
170 repeated meteorology and atmospheric  $CO_2$  concentration from 1996-2006. The current  
171 climate (i.e., no climate change ‘NoCC’) is considered the climate change scenario  
172 comparison baseline. At the start of the simulations, we created a Composite Forest Matrix  
173 (CFM), following the approach described in Dalmonech et al. [15], to simulate the  
174 potential effect of climate stressors on stands of different ages. The 3D-CMCC-FEM has  
175 been run at each site to cover the rotation period of each species (from 1997 to 2099) amid  
176 the current climate scenario (fixed atmospheric  $CO_2$  concentration at the year 2000 of  
177  $368.8\ \mu mol\ mol^{-1}$ ) consisting of detrended and repeated cycles of the present-day observed  
178 meteorology from 1996 to 2006 and the Business-as-Usual (BAU) management practices  
179 observed at each site (see Reyer et al., 2020 for the description of BAU applied at each  
180 site). Data required to re-initialize the model at every tenth of the rotation length were  
181 retrieved from each simulation. Hence, ten additional stands were chosen for each age in  
182 the composite matrix and added to the CFM. This collection of virtual forest stands was  
183 used to set different starting stand ages at the present day ( $age_{t0}$ ) due, ideally, to the past  
184 silvicultural practice and climate. Under this framework, a landscape of eleven different  
185 stands (in age and their relative C-pools and forest structure) for each site is created. These  
186 new stands were used, each running from 2006 to 2099, to assess the impact of climate  
187 forcing, as the model has already been shown to be sensitive to forest stand development  
188 and the relative standing biomass.

189 The 3D-CMCC-FEM was initialized with the structural attributes of the newly created  
190 stands from 1997, which was the starting year of all simulations and for all stands.  
191 Modeled climate change simulations under different RCP-emissions scenarios started to

192 differentiate in 2006 (up to 2100). The simulation runs from the different stand initial  
193 conditions, corresponding to different  $age_{t0}$  classes, were carried out without forest  
194 management as we are interested in the direct climate impact on undisturbed forest stand  
195 response, avoiding the confounding effects of forest management on the responses (for  
196 forest management effects, see Dalmonech et al., 2022). 825 different simulations were  
197 performed as they combined 5 ESMs \* 5 climate scenarios (4 RCPs + 1 current no climate  
198 change scenario) \* 11  $age_{t0}$  classes \* 3 sites. Results are reported for MAI (Mean Annual  
199 Increment;  $m^3 ha^{-1} year^{-1}$ ) and TCWS (Total Carbon Woody Stocks;  $MgC ha^{-1}$ ),  
200 respectively, as they are considered some of the most representative and fundamental  
201 variables in the carbon cycle and forestry. Following the methodology reported by  
202 Dalmonech et al. [15] (see Table S1 in Supplementary Materials), we evaluated the model  
203 forced with the modeled climate. We compared GPP and NPP<sub>woody</sub> against eddy covariance  
204 estimates and ancillary data for the years 1997-2005 for DK-Sor and FI-Hyy and 2000-  
205 2005 for CZ-BK1. We also compared the diameter at breast height (DBH) in all sites with  
206 field measures (see Supplementary Materials).

### 208 **3. Results**

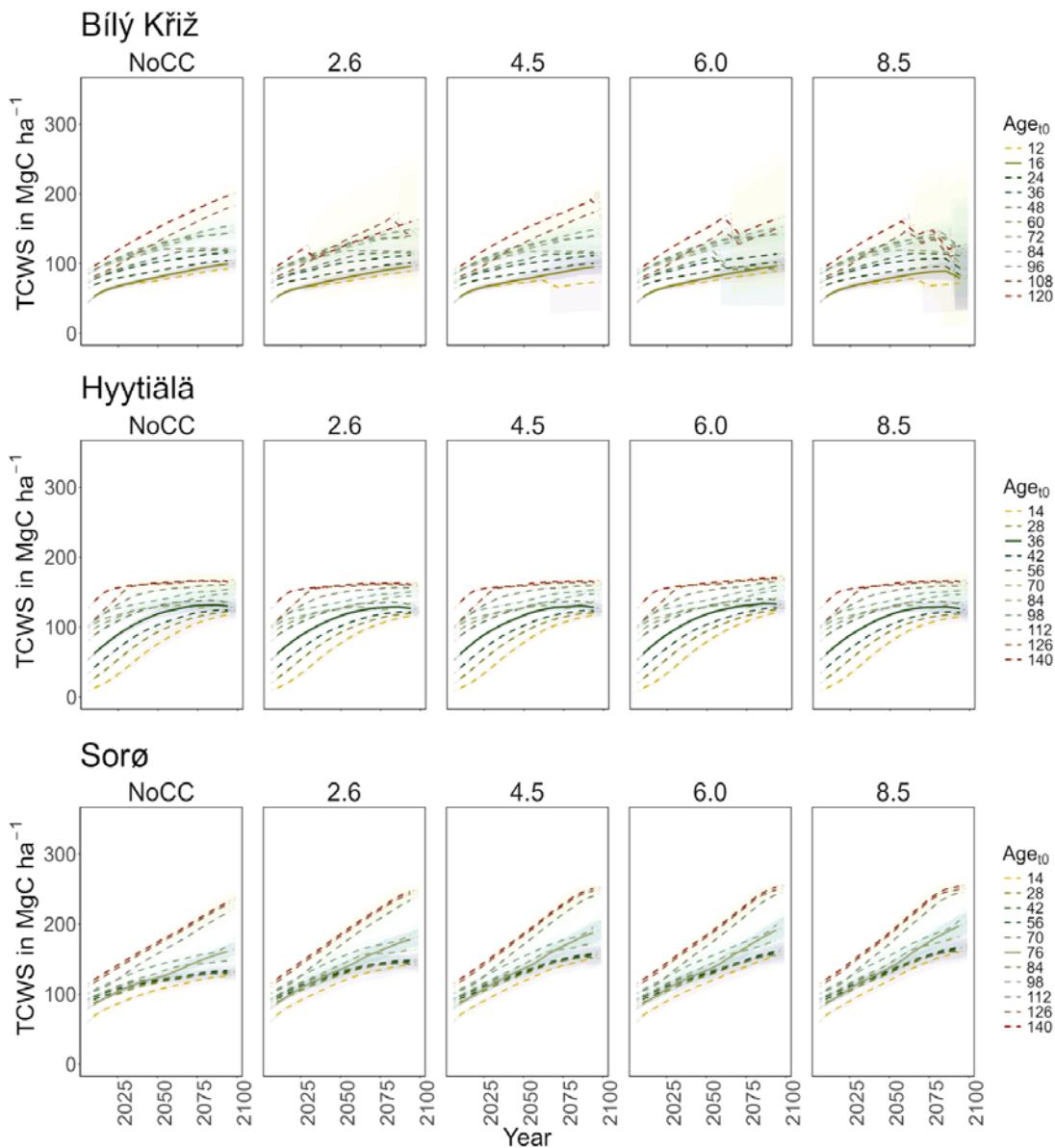
#### 209 *3.1 Effect of age classes and climate change on total carbon woody stock and increments*

210 Norway spruce at CZ-BK1 shows mean TCWS values ranging between ~70 to ~140  $MgC$   
211  $ha^{-1}$  under the NoCC scenario over the century, while from ~70 to ~130  $MgC ha^{-1}$ , with a  
212 decreasing pattern across all RCPs (Figure 2). In the Norway spruce stands under some  
213 ESMs climate forcing (HadGEM2-ES and GFDL-ESM 2M mostly) and under all climate  
214 change scenarios, the 3D-CMCC-FEM simulates mortality events for carbon starvation,  
215 which increase across stands under gradually warmer climate scenarios and from the oldest  
216 stands to the progressively youngest ones.

217 Under RCP 8.5, all classes show signs of decay at the end of the century. In the youngest  
218  $age_{t0}$  classes, a sharp decrease in MAI was observed (from 8 to 4  $m^3 ha^{-1} year^{-1}$ ), while in  
219 the older ones, it holds steady to ~3  $m^3 ha^{-1} year^{-1}$  with a peak around 2075 (Figure 3). At  
220 FI-Hyy, younger  $age_{t0}$  classes (14- to 42-year-old) showed the fastest increase in TCWS  
221 (reaching 120-130  $MgC ha^{-1}$  at the end of the century under all scenarios), also reflected in  
222 the pattern of MAI. Older  $age_{t0}$  classes showed a more stable trend throughout the  
223 simulation (Figure 2), culminating at ~150  $MgC ha^{-1}$ , with MAI steadily declining from 2.5  
224 to 2  $m^3 ha^{-1} year^{-1}$ . In all scenarios, the Scots pine peaked in the 126 and 56  $age_{t0}$  in TCWS  
225 and MAI, respectively. Minor differences were found in mean TCWS between the NoCC

226 and other RCP scenarios, ranging from  $-1.6\%$  (140-year-old class under RCP 2.6) to  
227  $+2.8\%$  (14-year-old class under RCP 6.0). At DK-Sor, results for TCWS show different  
228 patterns to other sites, with the highest values ranging between  $\sim 240 \text{ MgC ha}^{-1}$  (under  
229 NoCC) to  $\sim 255 \text{ MgC ha}^{-1}$  (under RCP 8.5) at the end of the century with the least TCWS  
230 under NoCC. The younger classes showed a shallow increase in TCWS during the  
231 simulation period, stabilizing at the end of the century, while the older ones kept growing  
232 (Figure 4). DK-Sor was the only site where the tightening of the climate conditions caused  
233 a positive effect on the MAI, particularly in the younger classes, reversing the trend from  
234 negative to positive at the end of the century.

235 In summary, a positive growth trend of TCWS over time was found in all sites, with the  
236 oldest  $\text{age}_{t0}$  classes accounting for the most carbon accumulation. Both conifer stands show  
237 a plateau with a reduction in growth at the end of the simulation, which is more  
238 pronounced and more severe in the warmest climate scenario. Conversely, the beech stands  
239 show a positive growth pattern in all scenarios. Similar results were obtained for MAI,  
240 where the conifers showed a decreasing trend over the simulation period despite different  
241 magnitudes and patterns among  $\text{age}_{t0}$  classes. The beech stands exhibited smaller variations  
242 among  $\text{age}_{t0}$  than among scenarios concerning other sites.



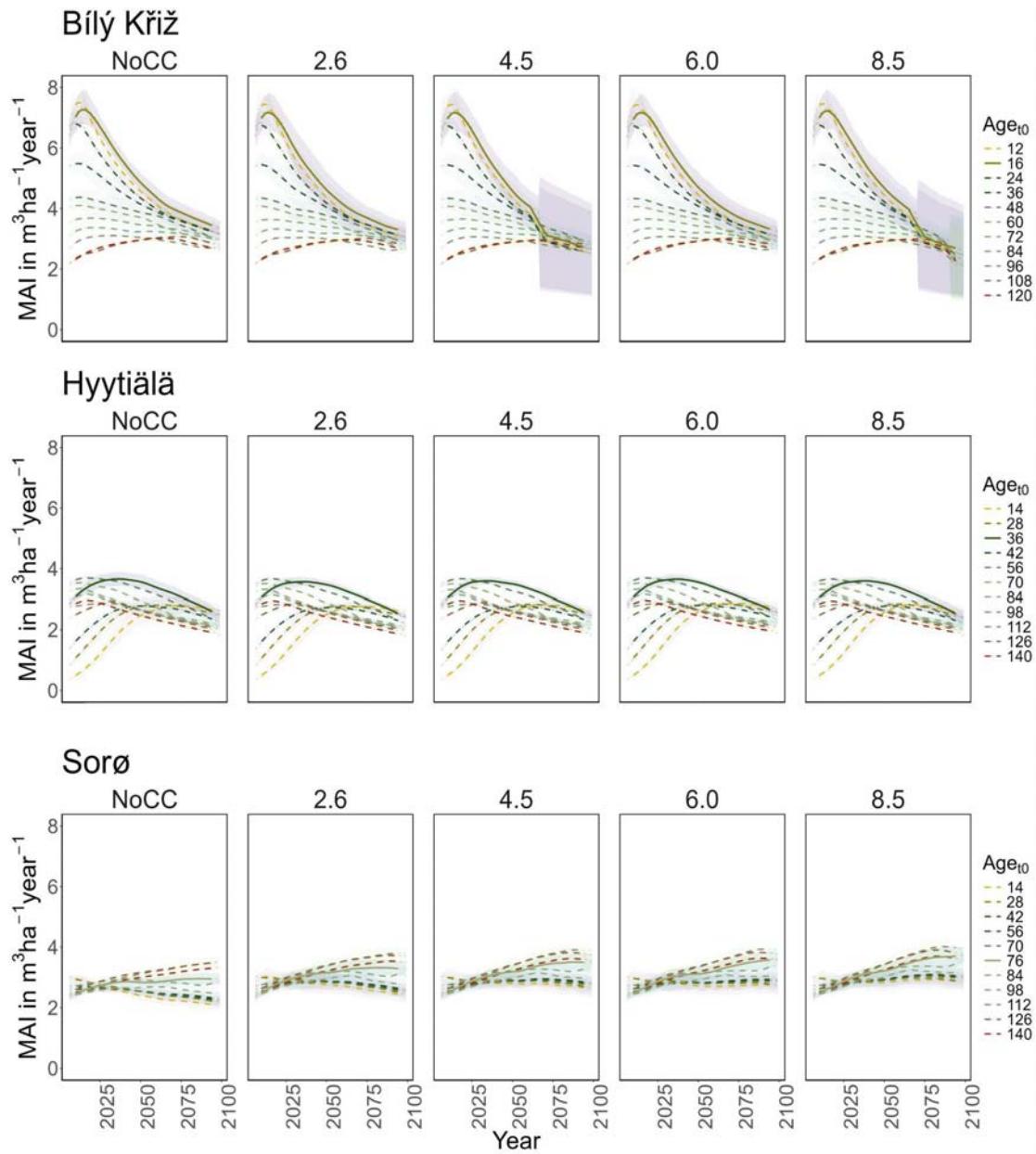
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245 **Figure 2.** Modeled TCWS ( $\text{MgC ha}^{-1}$ ) for age classes at the three sites in all scenarios along the simulation period (2006-2099). Lines represent the moving average of 10 years. The solid line corresponds to the real stand, while the dotted lines correspond to the virtual ones. The shaded area represents two standard deviations from the mean predictions with the results from the five ESMs' climate change scenarios.

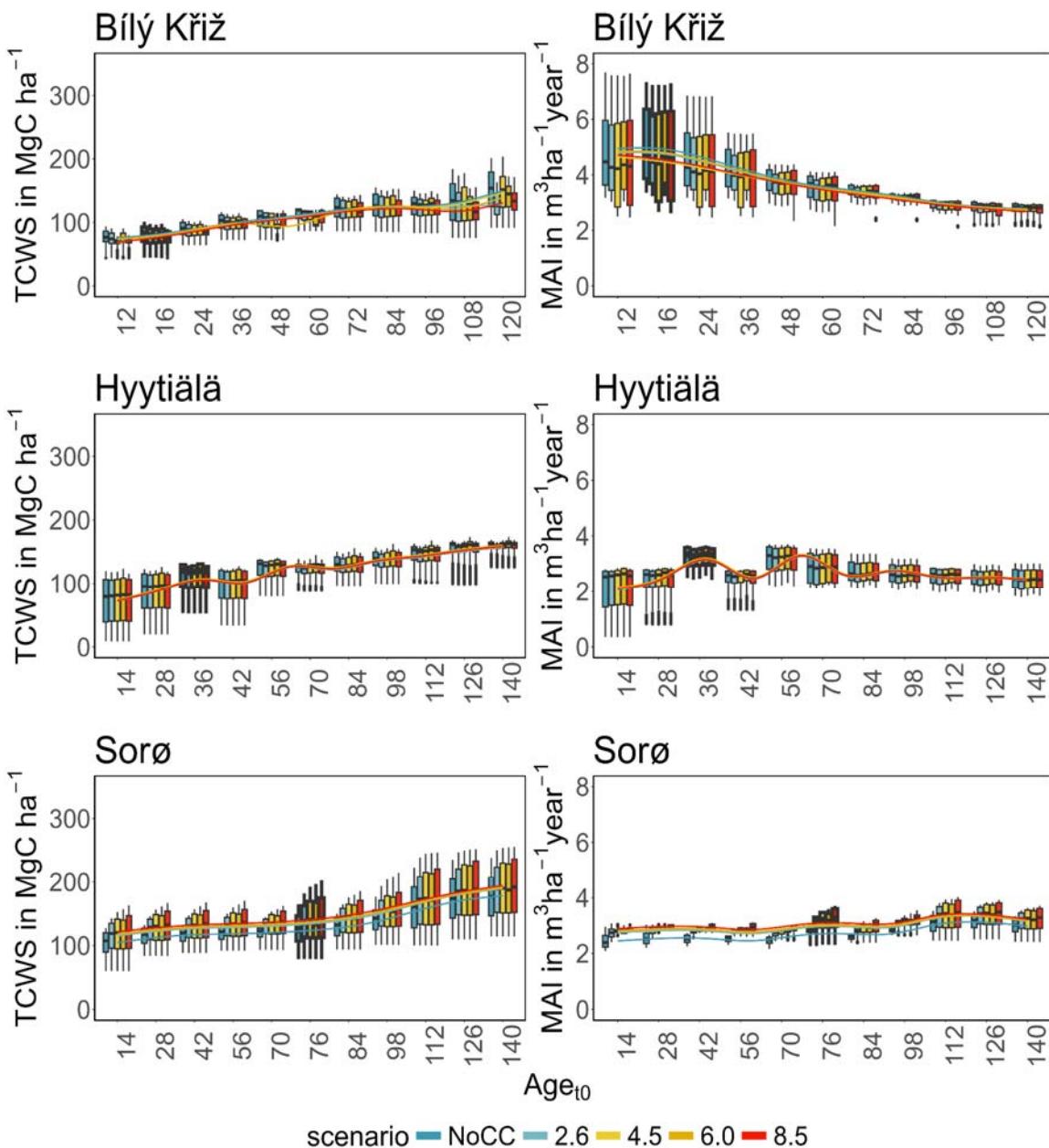
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Figure 3. Modeled MAI ( $\text{m}^3 \text{ ha}^{-1} \text{ year}^{-1}$ ) for age classes at the three sites in all scenarios along the simulation period (2006-2099). Lines represent the moving average of 10 years. The solid line corresponds to the real stand, while the dotted lines correspond to the virtual ones. The shaded area represents two standard deviations from the mean predictions with the results from the five ESMs climate change scenarios.



253  
254 **Figure 4.** Boxplot of modeled TCWS (left,  $\text{MgC ha}^{-1}$ ) and MAI (right,  $\text{m}^3 \text{ha}^{-1} \text{year}^{-1}$ ) for age classes at the  
255 three sites in the four RCPs scenarios compared to the NoCC (No Climate Change). Boxplots with thick  
256 borders correspond to the real stand. Lines are fitted throughout the median of the values of the variables  
257 using a generalized additive model.

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**Table 1.** Mean values of TCWS and MAI over the simulation period (2006-2099) for each scenario and age class

		Scenario				Scenario				
		N	R C P	R C P	R C P	N	R C P	R C P	R C P	
		o	2	4	6	8	o	2	4	
		C	.	.	.	.	C	.	.	
		C	6	5	0	5	C	6	5	
		A								
		Age					TCWS (MgCha <sup>-1</sup> )			
		1	7 2 .	6 9 .	6 5 .	7 0 .	6 5 .	4 4 .	4 4 .	4 4 .
		2	2 2 .	9 2 .	0 4 .	4 2 .	9 3 .	8 7 .	7 5 .	8 0 .
		1	7 .	7 3 .	7 3 .	7 4 .	7 1 .	5 .	4 .	4 .
		6	1 .	6 0 .	2 0 .	1 0 .	8 8 .	0 3 .	8 6 .	9 7 .
		1	8 6 .	8 4 .	8 4 .	8 5 .	8 2 .	4 .	4 .	4 .
		2	5 .	4 .	0 .	0 .	7 0 .	7 3 .	6 0 .	4 3 .
		4	3 .	7 .	6 .	2 .	0 .	0 .	1 .	3 5 .
		3	9 7 .	9 5 .	9 5 .	9 6 .	9 3 .	4 .	4 .	4 .
		6	4 .	3 7 .	0 0 .	1 3 .	6 4 .	3 0 .	1 9 .	2 1 .
		C	1 0 .	9 9 .	9 9 .	9 1 .	9 9 .	3 3 .	3 3 .	3 3 .
		Z	1 .	9 .	9 .	1 .	9 .	3 .	3 .	3 .
		B	4 .	7 6 .	3 0 .	0 .	5 8 .	8 1 .	7 1 .	7 4 .
		K	1 8 .	6 9 .	3 0 .	0 .	0 8 .	1 8 .	7 1 .	7 4 .
		1	1 1 .	1 0 .	1 0 .	1 9 .	1 0 .	3 3 .	3 3 .	3 3 .
		6	0 .	8 .	7 .	9 .	7 .	3 .	3 .	3 .
		0	1 .	1 .	8 .	3 .	0 .	5 5 .	5 5 .	5 4 .
		6	9 .	7 .	8 .	4 .	0 .	8 1 .	1 0 .	3 7 .
		0	1 .	1 .	1 .	1 .	1 .	3 .	3 .	3 .
		7	1 .	1 .	1 .	1 .	1 .	3 .	3 .	3 .
		2	9 .	0 .	6 .	0 .	5 .	4 2 .	3 5 .	3 4 .
		7	2 .	1 .	4 .	6 .	5 .	1 2 .	1 5 .	1 7 .
		2	1 .	1 .	1 .	1 .	1 .	3 1 .	3 1 .	3 1 .
		4	1 .	1 .	9 .	9 .	0 .	1 .	1 .	0 .

		.	.	.	.	8	2	1	4	8
		5	3	0	5	5				
		6	8	4	8	0				
		1	1	1	1	1				
		2	1	1	1	1				
		0	8	7	9	6	2	2	2	2
		.	.	.	.	.	.	.	.	.
		9	1	1	9	7	9	8	8	8
		6	9	2	8	0	0	8	3	0
		1	1	1	1	1				
		2	2	2	1	1				
		6	1	4	7	2	2	2	2	2
		1	.	.	.	.	.	.	.	.
		0	7	2	2	5	3	7	7	6
		8	8	6	8	7	3	4	0	1
		1	1	1	1	1				
		4	2	4	3	2				
		5	4	2	5	7	2	2	2	2
		1	.	.	.	.	.	.	.	.
		2	5	4	0	8	7	6	6	6
		0	0	3	1	2	4	2	9	5
		6	6	6	6	6				
		6	6	7	8	6	1	1	1	1
		1	.	.	.	.	.	.	.	.
		4	6	3	4	9	9	9	9	9
		2	2	9	9	8	1	3	5	4
		7	7	7	8	7				
		9	9	9	1	9	2	2	2	2
		.	.	.	.	.	.	.	.	.
		2	8	0	7	0	3	2	1	2
		8	3	0	3	5	7	1	9	1
		3	0	2	3	4	2	3	3	3
		1	0	0	0	0	0	1	1	1
		4	2	3	4	2	3	3	3	3
		3	.	.	.	.	.	.	.	.
		6	3	4	1	7	8	2	1	2
		4	2	2	7	5	6	0	4	1
		8	8	8	9	8				
		9	8	9	0	9	2	2	2	2
		.	.	.	.	.	.	.	.	.
		4	8	5	3	7	0	3	3	3
		2	2	5	9	8	5	1	3	7
		1	1	1	1	1				
		1	1	1	1	1				
		8	6	7	9	7	3	3	3	3
		.	.	.	.	.	.	.	.	.
		5	7	6	5	2	3	2	1	2
		6	6	2	7	0	6	0	4	7
		1	1	1	1	1				
		1	1	1	2	1				
		9	7	8	0	8	2	2	2	2
		.	.	.	.	.	.	.	.	.
		7	2	4	6	2	6	9	8	9
		0	9	3	1	5	5	1	6	9
		8	1	1	1	1	1	2	2	2

	4	2	2	2	2	2	.	7	7	.	7	.
		5	3	5	6	5	7	6	2	4	7	4
		.	.	.	.	.	6	2	.	.	7	4
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	8	.	.	.	.	.	6	6	6	6	6	6
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	9	9	1	1	1	1	2	2	2	2	2	2
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	4	3	9	8	4	2	6	2	9	7	7	6
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	4	4	6	8	8	3	0	6	2	0	0	0
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		3	2	5	4	8	2	2	2	2	2	2
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	2	5	6	3	8	1	5	7	8	8	8	9
	4	6	8	1	2	6	4	8	5	3	3	2
	1	1	1	1	1	1	.	2	2	2	2	2
		1	2	2	2	2	.	2	2	2	2	2
		5	5	7	7	0	2	2	2	2	2	2
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	6	9	0	7	2	5	5	7	8	8	8	8
	7	7	7	3	3	8	2	6	2	0	0	9
	0	1	2	2	2	3	3	4	5	3	3	2

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		1	1	7	3	6				
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		6	5	5	6	5	6	0	6	4
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		4	3	6	6	9	2	2	2	2
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		8	6	9	6	1	5	5	7	8
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		8	7	1	6	9	3	2	7	6
		1	1	1	1	1				
		5	6	6	6	6				
		4	4	6	5	8	3	3	3	3
		.	.	.	.	.	.	.	.	.
		1	1	3	5	7	0	2	2	2
		1	2	2	8	1	7	0	1	5
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		6	7	7	7	8				
		6	6	8	7	0	3	3	3	3
		.	.	.	.	.	.	.	.	.
		1	2	3	3	6	3	0	2	2
		6	4	3	8	3	2	6	4	8
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		7	8	8	8	8				
		0	0	1	1	4	2	3	3	3
		.	.	.	.	.	.	.	.	.
		1	4	0	8	2	0	9	0	0
		0	9	8	4	6	5	1	8	1

265

#### 266 4. Discussion

##### 267 4.1 Age-dependent impacts of climate change on forests' increment and C-stocks

268 The successional stage, represented by forest age, was the main driver controlling C-  
 269 storage capacity and biomass accumulation, as already known by previous studies [48, 49,  
 270 50], with differences greater among different age cohorts under the same scenario than in  
 271 different climate scenarios under the same age class [13, 15]. The evidence that the carbon  
 272 budget is mainly controlled by stand age suggests that the effects of climate change on  
 273 forest cohorts are generally less significant than the effect of age, mainly in terms of the

amount of standing biomass. In this sense, age represents multiple and interacting processes, such as tree size [51, 52], forest structural traits (canopy closure and LAI), reduction in stomatal conductance [17], and adaptation to specific environmental conditions which, in turn, make it possible to increases the above-ground biomass (AGB) [53]. The model could reproduce the expected behavior of biomass (and thus carbon) accumulation, simulating rapid growth at a young age and saturation for the oldest age class, but not necessarily at the end of the simulation period. Approaching the physiological optima for the species may benefit the biomass synthesis through an augmented photosynthate supply but may eventually increase the respiratory costs of tissue growth and maintenance despite a strong acclimation capacity [19]. High respiratory costs in warm climates and with low precipitation regimes combined in the older age classes lead to C-starvation and mortality phenomena, as observed for the Norway spruce at the CZ-BK1 site. This indicates that the environment has reached its carrying capacity and that competition for limited resources, such as light and water, is excessively high to sustain more biomass in the oldest age classes.

We found different C-accumulation patterns under climate change between coniferous stands and broadleaves. As expected, increasing temperature and changes in precipitation patterns led to a decline in above-ground biomass in spruce stands, especially in the older age classes. On the contrary, the results show that beech forests at DK-Sor will maintain and even increase C-storage rates under most RCP scenarios. Scots pine forests show an intermediate behavior with a stable stock capacity over time and in different scenarios but with decreasing MAI. These results confirm current observations worldwide that indicate a stronger climate-related decline in conifers forests than in broadleaves [54, 55, 56, 57]. This contrasting response is explained by the different characteristics of the two *phyla*, in particular, due to the temperature adaptation, with generally lower optimum temperature in conifer and less sensitivity to the length of the growing season. Similarly, conifers also show lower efficiency in water management because of the shallower root system, which increases the sensitivity to soil aridity and its vulnerability to drought events [58]. Recent studies confirm that growth decline is more pronounced in conifers than broadleaf, especially beech forests, in the most northern species distribution [59]. Our results confirm the same growth patterns found by recent studies [50, 57, 60], where broadleaves outperform conifers in productivity, and climate warming will probably exacerbate these opposite growth patterns.

307 However, despite some studies suggesting that age modulates different adaptation  
308 strategies to some extent, it remains unclear whether younger trees may be more affected  
309 by climate change than older ones. Bennett et al. [61], in a global analysis, found that  
310 droughts consistently had more severe impacts on larger (older) trees, while Wang et al.  
311 [11] observed a more substantial and sharper decline in basal area increment in young  
312 Korean pine in China. Hogg et al. [62] found that the percentage decrease in biomass  
313 growth was not significantly different for young, productive stands compared to older, less  
314 productive ones. Our study suggests that warmer and drier conditions and extended  
315 growing seasons will affect younger stands more than older ones, but with different trends  
316 among species. In particular, MAI will be positively affected in younger beech forests,  
317 while it will remain stable in older stands. On the contrary, climate change will strongly  
318 impact the growth rate of young conifers stands more than older ones. Older forests tend to  
319 be more stable and resilient than younger ones due to their rugged and stable interaction  
320 with climate triggers and better responsiveness to environmental changes. The year-to-year  
321 climate variability is buffered by larger carbon pools in sapwood and reservoirs in older  
322 trees, leading to higher long-term stability than younger trees [13]. In this sense, ages  
323 represent the “memory” of the forest to past climate and disturbance regimes, which align  
324 the species-specific traits to the environmental conditions in which they grow, creating the  
325 niches in which AGB accumulates [53, 63].

326 Despite numerous efforts to decipher forests' response to climate change, the intricate  
327 methods employed by tree species to withstand extreme climates still need to be fully  
328 unveiled. Further research exploiting ecophysiological models explicitly accounting for  
329 age, tree-ring experiments, and remote sensing will be critical to understanding forest  
330 ecosystems' adaptation strategies to climate change, particularly in the face of rapid  
331 warming and extreme disturbances. A better understanding of the interaction between  
332 forests and climate can inform better forest management strategies, ultimately dampening  
333 the impacts of climate change on forest ecosystems.

## 334 335 **5. Limitations**

336 The presented modeling framework has some limitations that should be considered. Firstly,  
337 natural disturbances as consequences of climate change, such as windstorms, forest fires,  
338 and insect outbreaks, were not simulated. In contrast, climate extreme events are  
339 considered to be already included in the climate scenarios used to force the model and,  
340 thus, already accounted for in the model outputs. Additionally, other indirect alterations

341 due to climate change of key drivers, such as nitrogen deposition, phosphorus, or ozone,  
342 which can somewhat amplify or reduce our results, were not assessed. Nonetheless, some  
343 studies (e.g., [64]) lend credence to the notion that this phenomenon may not be applicable  
344 across the board. They highlight the significant responsiveness of various tree species to  
345 CO<sub>2</sub> fertilization across a wide range of nutrient availability. Finally, no allowance was  
346 made for the possibility of species migration to and from the study areas. However, these  
347 dynamics may require longer timescales than those simulated in this study.

348 **5. Conclusions**

349 Forest age is confirmed to be a significant factor in determining the carbon storage capacity  
350 and biomass accumulation in forest ecosystems, especially in the context of future climate  
351 uncertainty. The effects of species, site location, stand-level characteristics, and  
352 development stage vary significantly and are contingent on specific factors. We observed  
353 that differences in biomass accumulation were more pronounced among different age  
354 cohorts than among different climate scenarios within the same age class, with contrasting  
355 carbon accumulation patterns under climate change between coniferous and broadleaf  
356 forests. Furthermore, our findings shed light on the differential impacts of climate change  
357 on younger versus older forest stands. Warmer and drier conditions are projected to affect  
358 younger stands more severely, particularly in coniferous forests. However, older forests  
359 will likely exhibit greater stability and resilience due to their accumulated carbon pools and  
360 enhanced adaptability to environmental changes. While our study provides valuable  
361 insights, it also underscores the need for further research to unravel the complex  
362 mechanisms by which forests adapt to climate change. This deeper understanding can  
363 inform more effective forest management strategies, helping to mitigate the impacts of  
364 climate change on forest ecosystems in the future. The varying responses of different tree  
365 species highlight the need for tailored management approaches and conservation efforts to  
366 enhance the resilience of our forests.

367  
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369 original draft, Writing – review & editing; D.D.: Data curation, Formal analysis,  
370 Investigation, Writing – review & editing; M.M.: Writing – review & editing; E.G.:  
371 Writing – review & editing; F.G.: Writing – review & editing; G.D.: Writing – review &  
372 editing; G.C.: Writing – review & editing; A.C.: Formal analysis, Investigation, Writing –  
373 original draft, Writing – review & editing, Conceptualization. All authors have read and  
374 agreed to the published version of the manuscript.

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379

380  
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## 642 Supplementary materials

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### 644 Model evaluation

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Daily model outputs were evaluated against eddy-covariance and measured structural data at the site level, in terms of percentage root mean squared error (RMSE%) and Fractional mean Bias (FMB). The GPP evaluation for simulations forced with observed site-specific daily weather data (1997-2005 for FI-Hyy, and DK-Sor, and 2000-2005 for CZ-BK1) resulted in an RMSE% of 1.05, 1.52, and 1.43, with r values of 0.92, 0.87 and 0.94 for FI-Hyy, CZ-BK1 and DK-Sor, respectively (Table 1). Similar results were obtained for NPP<sub>woody</sub> in the site of DK-Sor and CZ-BK1 ( $351 \pm 61 \text{ gC m}^{-2} \text{ year}^{-1}$  vs.  $346 \pm 36 \text{ gC m}^{-2} \text{ year}^{-1}$  measured, and  $442 \pm 79 \text{ gC m}^{-2} \text{ year}^{-1}$  vs.  $380 \pm 38 \text{ gC m}^{-2} \text{ year}^{-1}$  measured, respectively). At FI-Hyy, modeled NPP<sub>woody</sub> data was overestimated in respect to the measured values ( $317 \pm 21 \text{ gC m}^{-2} \text{ year}^{-1}$  vs.  $228 \pm 23 \text{ gC m}^{-2} \text{ year}^{-1}$  measured).

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**Table S1** | Performance statistics (relative root mean square error RMSE ( $\text{gC m}^{-2} \text{ day}^{-1}$ ) and Fractional Mean Bias, FMB) computed from monthly seasonal values and annual series of model gross primary productivity, GPP, against eddy covariance estimated and diametric annual increment data, DBH increment, against measured data. Results are reported for simulations forced with local and modeled climate (i.e., ESM) (ESM1, 2, 3, 4, 5 refer to HadGEM2-ES, IPSL-CM5A-LR, MIROC-ESM-CHEM, GFDL-ESM 2M, and NorESM1-M, respectively).

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		GPP				DBH increment			
		SEASONAL				DBH increment			
		CLI	R		F	R		F	R
			F	M		M	S		M
C Z - B K 1	Loca 1	ES	-		-			-	
			0	1	0			0	
			.	7.	.	18		.	0.
			0	9	0	0.		1	2
			6	2	6	96		7	6
	M1 ES	M1 ES	-		-			-	
			0	2	0			0	
			.	8.	.	16		.	0.
			2	6	0	1.		2	2
			2	8	7	16		9	4



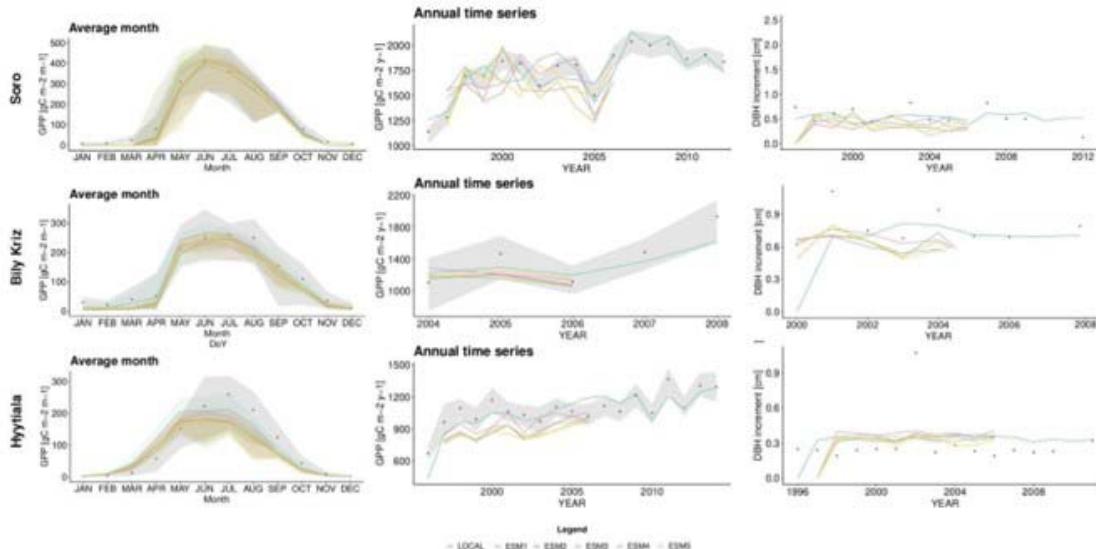
		2 4	6	2		1
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		.	9.	.	21	0
	ES	1	0	0	0.	6
	M1	2	1	7	27	2
		-	2	0	-	
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K	ES	.	6	0	3.	5
-	M2	7	3	2	82	1
S			-	-	-	
o		-	3	0	-	
r		0	3.	.	15	0
	ES	.	4	0	0.	0.
	M3	7	6	2	62	6
		-	-	-	-	
		0	2	0	-	
		.	6.	.	19	0
	ES	1	5	0	5.	6
	M4	2	5	7	54	1
		-	-	-	-	
		0	2	0	0	
		.	7.	.	16	0.
	ES	0	6	0	8.	5
	M5	9	9	5	62	9

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669 **Figure S1** Evaluation of monthly seasonal GPP ( $\text{gC m}^{-2} \text{ month}^{-1}$ ) fluxes (left column) and annual ( $\text{gC m}^{-2}$   
670 year<sup>-1</sup>) fluxes (central column) for the sites of Sorø, Bily Kriz, and Hyytiala (rows). Quality-checked and -  
671 filtered GPP values evaluated at the sites by the eddy covariance technique are reported as black dots. The  
672 shaded area for seasonal values reports the maximum and minimum monthly values recorded in the time  
673 series. The shaded area for annual data represents the relative uncertainty bounds. In the third column, a  
674 comparison of the predicted annual DBH increment (cm y<sup>-1</sup>) with site observations at the three sites is  
675 reported. Measured data are shown as black dots. Simulated data are reported as continuous lines.

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